

Erosion and accumulation flood landforms in Sázava River in spring 2006

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Abstract

This article deals with channel bars, flood accumulations of floodplain and bank scours as the most important geomorphologic effects of the spring flood in Sázava River in 2006. The article deals particularly with its spatial distribution with regard to the channel position and weirs. Relation between basic characteristics of sediments of these channel bars is surveyed. Spatial and statistical analyses were done in ArcGIS 9.2 and STATISTICA 8 Standard based on input data from field geomorphological mapping after the flood in 2006, it follows that there is a relation between spatial distribution of channel bars and flood accumulations in floodplain. There is also a relation between channel bars and bank scours. It emerged that spatial distribution of channel bars, flood accumulations in floodplain and bank scours is uneven and connected to weirs whose immediate effective sedimentary influence reaches in Sázava River to 60 m distance. Smaller channel bars connected to banks are typical for parts of river which are not influenced by weirs. Channel bars can be used for placement prediction of flood accumulations in floodplain.

Key words: floodplain – flood accumulations – channel bars – bank scours – weirs – flood – Sázava River

Introduction

Floodplain is an exposed area where anthropogenic and fluvial (natural or anthropogenic influenced or accelerated) geomorphologic processes meet (e.g. Whittow, 1984; Collin, 1988; Huggett, 2003; Brierly, Fryirs, 2005; Křížek, 2007a). Overbank floods present decisive processes for channel formation (Hickin, 1983). Monitoring of erosion and accumulation fluvial processes in spatial (within and outside the channel) and time context enables prediction and insight of channel-floodplain geosystems' functioning (sensu Knighton, 1998; Lehotský, 2005). Floodplain sediments are controls on channel behavior (Schumm, Thorne, 1989). Effects of fluvial accumulation and erosion in channels and floodplains are mainly studied only in selected parts of rivers (e.g. Walling, He, 1998; Steiger, Gurnell, 2001; Hooke, 2003; Pierce, King, 2008). Discrete approach to research of fluvial accumulation and erosion effects together with change of type of channel bars depending on local channel morphology (sensu Zielinski, 2003) limits an evaluation of continual spatial dependences of channel bars and overbank accumulations (in longitudinal profile). Nicholas and Walling (1997), Walling and He (1998), Steiger and Gurnell (2001) dealt with prediction of flood sediment deposition in floodplains (rather in their cross-

profiles) based on modeling of mean annual deposition rates. Czech geomorphology works which deal with geomorphologic effects of recent great floods (in 1997, 2002, 2006) concentrate more or less on the flood geomorphology effect inventory or they studied relation of the floods to anthropogenic channels' changes (Hrádek, 1999; Hrádek, 2005; Křížek, Engel, 2006).

The aim of this article is to characterize basic positional and qualitative characteristics of fluvial accumulation and erosion landforms which were created or were transformed during the flood in 2006 and to determine their mutual relations.

During springtime in 2006 snow melting on large areas in middle altitudes occurred which led to overflow of large number of rivers and resulting floods. Snow water storage was in total 446.3 mil. m³ in the whole catchment of Sázava River on 20th March 2006 (Šercl et al., 2006). These spring floods caused bankfull of Sázava River accompanied by creation of fluvial erosion and accumulation landforms in its floodplain and its channel. The two-top flood culminated on 30th March 2006 and maximum discharge of Sázava River was 536 m³/s in Nespeky, for comparison the mean annual discharge of Sázava River is 23.4 m³/s in Nespeky (Šercl et al., 2006).

The studied area, where the analysis of geomorphologic effects of the spring flood was elaborated, is formed by floodplain of Sázava River from its confluence with Losenický potok Brook in the upper part up to Pikovice where damming water of Vltava's cascade comes through (Fig. 1). A 183 km long section of floodplain was mapped during 27 days.

Sázava River is single-channeled, meandering stream with an irregular sinuosity. The valley of Sázava River is in the major part of the flow characteristic by its deep valley where difference in height of the valley bottom and its edge reaches up to

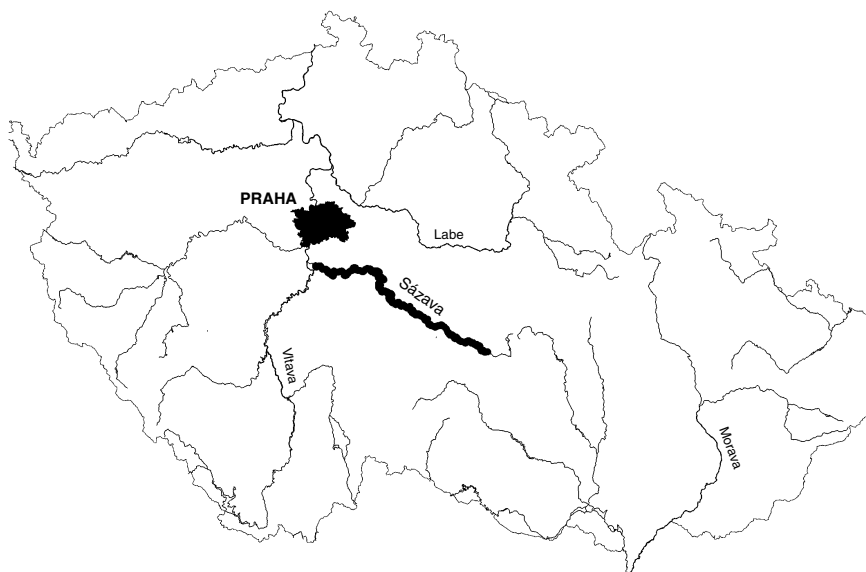


Fig. 1: Position of study area – Sázava River

200 m. According to various authors (Balatka, Sládek 1962) approximately 4 alluvial terrace levels in convergent direction towards the flow exist here. The lowest and youngest terrace level is located 2–10 m above the valley bottom and the highest and oldest terrace level is situated in highs 72–126 m (Balatka, Sládek, 1962).

From the geological viewpoint the major part of Sázava River valley is developed on basement composed of granites and granodiorites of Middle-Bohemian pluton which are locally crossed by dikes of paleovolcanic and metamorphic rocks. Lithologic situation reflects into sandy sediments (as a result of granite weathering) which were transported and accumulated by Sázava River during the flood.

Methods

Geomorphologic mapping of geomorphologic effects of the spring flood in 2006 was executed in period of time from 7 May to 4 July 2006 and focused on floodplain (actually on bottom of Sázava River valley) and landforms which extend into it from surrounding valley slopes. Basic maps of the Czech Republic with scale 1 : 10 000 (ČÚZAK) were basic documents for field mapping. The field mapping was executed according to unified legend created by M. Křížek (Křížek, 2006). Constituent parts of Sázava River were mapped and digitalized by 8 employees and students of Department of Physical Geography and Geoecology of Faculty of Science of the Charles University in Prague. The results obtained by the field workers were revised and completed by the author of this article; morphometrical and positional characteristics were set by GPS and laser range finder and clinometer MDL LaserAce-Hypsometer.

Tab. 1: Surveyed characteristics of selected landforms

Landform	Attributes
<i>Alluvial fan</i>	Thickness: max. thickness (height difference between foot and top of this landform) of its sediment in the floodplain
<i>Alluvial terrace</i>	Height: height above the current floodplain
<i>Floodplain /borders/</i>	Type: accumulation (1), transport (2) and erosive part (3) refers to floodplain only not the channel
<i>Flood accumulation in the floodplain</i>	Type: sandy (1), gravel (2), combined (3) + Origin: date of origin if known, e.g. 2002, otherwise 0
<i>Bank scours</i>	Origin: date of origin if known, e.g. 2002, otherwise 0
<i>Weir</i>	Height: up to 1 m (1); 1–2 m (2); above 2 m (3)
<i>Bridge, footbridge</i>	Length + Height above the bottom of the river-channel (maximal) + State: not damaged (0), damaged (1), destroyed (2)
<i>Bar</i>	Width of the river-channel: width of the channel in the locality of a bar + Location: (1–5) see fig. 1 + Length + Width + Connection: without connection to bank (0), connection to bank (1) + Material: sandy (1), gravel (2), combined (3) + Gradation: without gradation (0), normal gradation (1), inverse gradation (2) + Vegetation: without vegetation (0), grass (1), shrubs or trees (2)

The floodplain was defined in accordance with geomorphologic definition as flat accumulation plain along watercourse built from fluvial sediments; during floods it gets usually partly or completely flooded and its borders were determined based on its different morphology (inclination of slopes) from valley slopes and other landforms (*sensu* Křížek, Hartvich, Chuman et al., 2006). These landforms and buildings located in floodplain and its surrounding were classified: alluvial fans, fluvial (alluvial) terraces, fluvial (flood) overbank accumulations in floodplain, bank scours, weirs, bridges and channel bars (Tab. 1; Křížek, 2006). Channel bars represent fluvial sediments in the river channel. Width of the river channel was measured for every bar (for corresponding part of channel where bar is situated). Its position within the channel was also determined. The channel was divided into 5 sections with the following cross-section profile orientation: section 1 lies near the left bank and section 5 lies near the right bank. In case a bar lies in more sections it was marked by rising numeral order of particular sections (e.g. 234). The channel bar length describes the size which is parallel to the water flow direction and the width is perpendicular to it. Other bars' feature was connection with banks (whether it is /e.g. lateral bar – *sensu* Brierley, Fryirs 2005/ or it is not connected with a bank). Psephite to psammite ratio was detected for each fluvial accumulation. From the grain size viewpoint the fluvial sediments were distinguished: gravel-cobbles (contain more than 25% of coarse grained /psephite/ clasts), sands (contain less than 10% of coarse grained /psephite/ clasts) and combined (contain 10–25% of coarse grained /psephite/ clasts) (*sensu* Kukal, 1985). Gradation means sediments segregation and their sorting by size: normal gradation means that the fluvial material which builds the channel bar gets more fine-grained sediment in the downstream direction; inverse gradation means that fine-grained sediments are deposited first than coarse sediments (*sensu* Bridge, 2003); channel bars without gradation do not have apparent sorting. The last surveyed feature was channel bar cover by vegetation: vegetation free, partly covered by grass, with shrubs or trees. This characteristic describes stability, actually duration of existence of the channel bar.

Spatial information about landforms and their other characteristics were processed in ArcGIS 9.2 (ESRI 2003) with the use of extension – Spatial Analysis and then were statistically analyzed (one-way ANOVA) in program STATISTICA 8 Standard (StatSoft 2003). Densities of channel bars, flood overbank accumulations and bank scours were determined in GIS for cell size 100 m and radius 1000 m. Distance relation between single landforms were determined with the use of buffering with 5 m step and conversion into the grid for cell size 5 m. The first distances category of flood overbank accumulations from the channel (i.e. 0–25 m) was determined regarding to changes of channel width (from 7 m on the upper stream to 75 m on the lower stream of Sázava River). Centroids of flood accumulations substitute real areas of these landforms for mutual position relations between channel bars and flood overbank accumulations in the floodplain. This was done in order to limit error of multiple-choice matching of larger flood accumulations to more than one channel bar. For the same reason centroid was used to calculate density of flood overbank accumulations in the floodplain.

Results

Floodplain and flood accumulation landforms

Floodplain of Sázava River is separated from other part of relief (e.g. from valley slopes or step of alluvial terrace) by an edge with more significant inclination change which appears in cross-section profile. Whereas surface of Sázava River floodplain has inclination to 2° , neighboring valley slopes have inclinations in range $7\text{--}90^\circ$, which is due to erosion valley of Sázava River cutting into the basement. In areas where alluvial fans of adjacent valleys reach floodplain of Sázava River the channel is being pressed away towards opposite bank which is usually accompanied by narrowing of floodplain

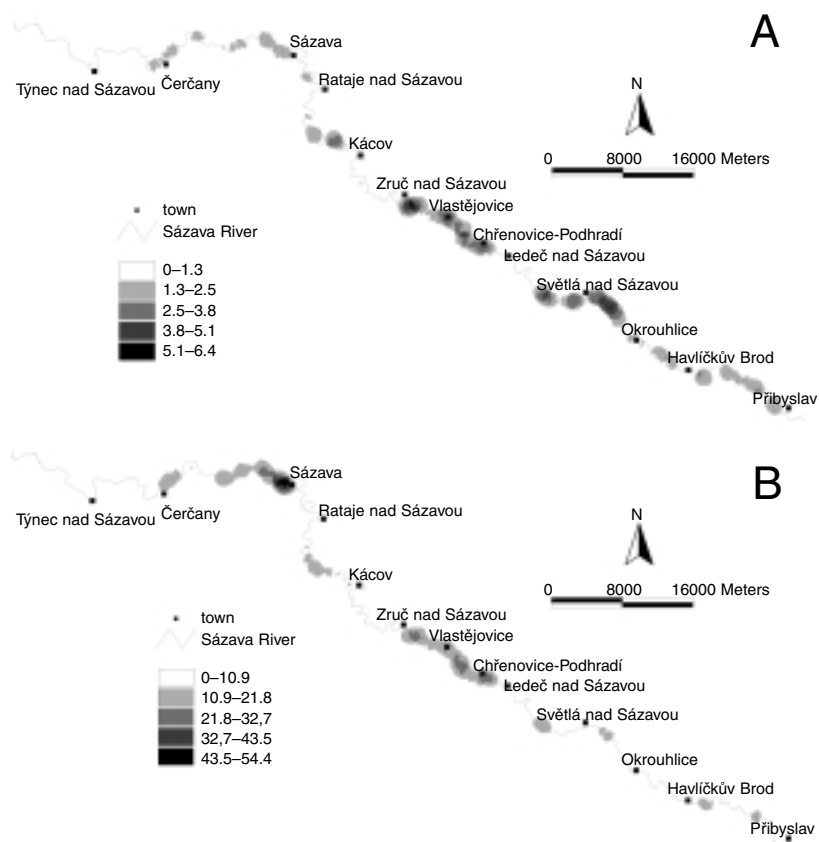


Fig. 2: Density of flood overbank accumulations in the floodplain. Density defines number of flood overbank accumulations per 1 km^2 for radius 1 000 m . A – Each flood overbank accumulation in the floodplain is represented by its centroid; thus the density is calculated with respect to centroid of each flood overbank accumulation in the floodplain. B – Each flood overbank accumulation in the floodplain is represented by several points. Number of these points depends on the area of relevant flood overbank accumulation in the floodplain; thus the density is calculated with respect to area of each flood overbank accumulation in the floodplain

width. Regarding the total length of Sázava River it can be stated that width of the studied part of floodplain is more or less constant, actually in the upper and middle part of Sázava River it increases very slowly in the direction of downstream. Further, in some areas in direction of downstream it gets more narrow (e.g. below Havlíčkův Brod) which is caused by geologic predisposition and also by anthropogenic impact into the floodplain, e.g. railway embankments. The average floodplain width is 115 m in the studied upper part of Sázava River, its range is from 20 to 389 m. The average width of floodplain in the studied middle part of Sázava River is 121 m, it ranges in interval from 42 to 432 m. Generally we can say that variation from the average is not as big as it is in the studied upper part of Sázava River. The average floodplain width is 110 m in the lower part of the studied floodplain of Sázava River, however the variation of widths of floodplain is the smallest, boundary values are 70–167 m.

During the flood in spring 2006 the floodplain was completely inundated in some places, inundation occurred even outside the floodplain in some localities. Flood sediments found on feet of alluvial fans or on lower terrace steps, which lie out of the floodplain, prove it. In the studied part of floodplain of Sázava River there were found 262 sandy flood overbank accumulations (86% from total number), their total area is 102 ha (79.4% from total area of all flood accumulations in the floodplain), 8 gravel-cobble flood overbank accumulations (3%), their area is 1 ha (0.7% from total area of all flood accumulations in the floodplain) and 28 combined flood overbank accumulations (9%), their area is 25 ha (19.4% from total area of all flood overbank accumulations) and 5 unrecognized flood accumulations (2%), their area is 0.7 ha (0.5% from total area of all flood accumulations in the floodplain). Mean thickness of these accumulations reaches 3–30 cm. Most of flood overbank accumulations can be found near Čerčany, below town Sázava, between Ledec nad Sázavou and Zruč nad Sá-

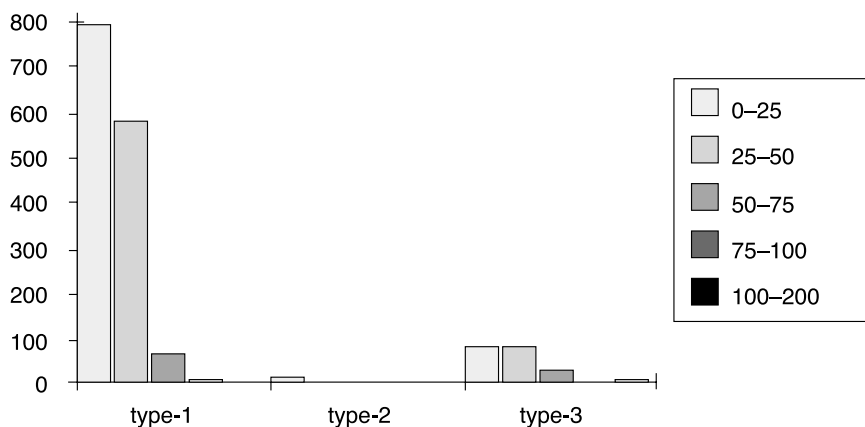


Fig. 3: Number of types of flood overbank accumulations regarding distance from the channel (midstream). Type 1 – sandy flood accumulations, 2 – gravel-cobble flood accumulations, 3 – combined flood accumulations. Category values (0–25, 25–50 etc.) represent distances from midstream in meters. Each flood overbank accumulation in the floodplain is represented by several points – with respect to area of each flood overbank accumulation in the floodplain

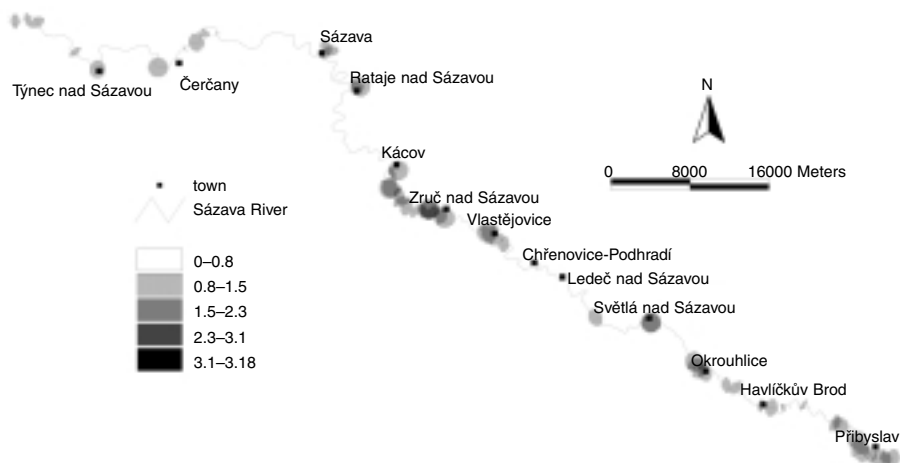


Fig. 4: Density of channel bars in Sázava River. Density represents number of channel bars per 1 km² for radius 1 000 m

zavou, surroundings Světlá nad Sázavou and below Příbyslav (Fig. 2). The same picture shows evident conformity between number of flood overbank accumulations in floodplain and their areas. Concentration of flood sediments is the highest in immediate vicinity of Sázava River channel (Fig. 3), distances up to 25 m from midstream of the channel prevail and flood overbank accumulations which are distant above 75 m from midstream of the channel occur sporadically. Combined type of flood overbank accumulations has the most frequent distance from midstream in range 40–60 m. It is related to multiphase process of accumulation in middle part of floodplain where these landforms are situated only in these positions (ranges). Most (73%) of flood accumulations can be found within 20 m from weirs.

147 channel bars were found in the studied part of floodplain in total, 27 (18%) were sandy, 81 (55%) gravel-cobble and 39 (27%) combined sediment. From the viewpoint of gradation 119 (81%) of channel bars without gradation were identified, 26 (18%) with normal gradation and 2 (1%) with inverse gradation. 82 (56%) of channel bars were not connected with banks so they formed islands while 65 (44%) were connected with banks. 64 (44%) channel bars had no vegetation cover, grass was found on 55 (37%) channel bars and shrubs or trees grew on 28 (19%) of them. The highest densities of channel bars were in surroundings of Příbyslav, Okrouhlice, Světlá nad Sázavou, Vlastějovice and then between Zruč nad Sázavou and Kácov, Rataje nad Sázavou, Sázava, Čerčany, Týnec nad Sázavou and Čísovice (Fig. 4).

Channel bars are located mainly in vicinity of weirs (Fig. 5, 6), most of them can be found within 20 m from a weir then their number falls rapidly. Channel bars with normal gradation occur mostly in immediate vicinity of weirs, within distance up to 20 m then their number smoothly falls. On the contrary, channel bars without gradation are mainly located in distance up to 20–60 m from weirs, their minimal number is located in distance of 60–100 m. Channel bars with inverse gradation were found only in distance of 60–100 m from weirs. From the viewpoint of channel bars' position in channel, they

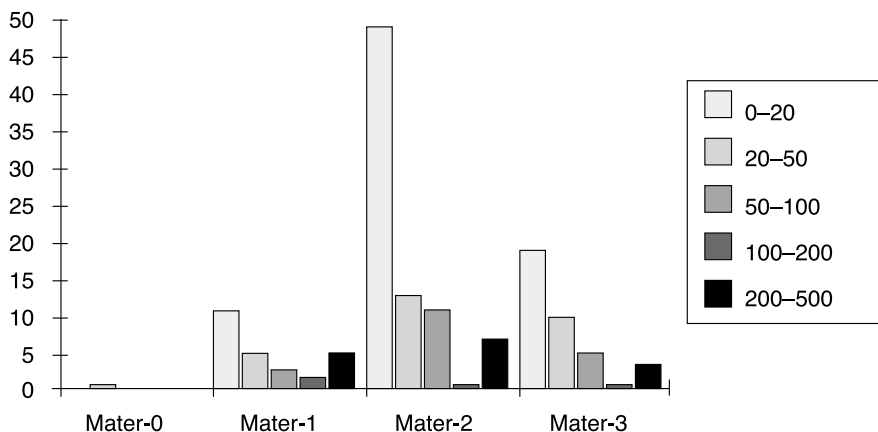


Fig. 5: Spatial distribution of channel bars regarding distance from weirs (in meters). Mater-0 ... nonspecified channel bars; Bars: Mater-1 ... sandy; Mater-2 ... gravel-cobble; Mater-3 ... combined

do not occur in channel sections 1 and 5 in immediate vicinity of weirs (up to 20 m), they occur the most in distance bigger than 100 m from weirs, on the contrary central positions (channel sections: 2, 3, 4, 23, 34, 234) of channel bars (e.g. medial bars or transverse bars – sensu Brierley, Fryirs, 2005) occur usually up to 60 m from weirs. Most



Fig. 6: Channel bar with grass cover bellow a weir near Chřenovice

of channel bars connected with banks (lateral bars – sensu Brierley, Fryirs, 2005) are located in distances bigger than 100 m from weirs. Majority of large channel bars with width more than 5 m are located in distance of 20–60 m from weirs, contrary small bars with width up to 3 m are spread out more regularly with maximal occurrence in distance above 100 m. Channel bars covered by grass, shrubs or trees or channel bars without vegetation cover can be found the most up to 60 m from weirs, only bars with shrubs and trees occur much less in immediate vicinity of weirs (up to 20 m). Sandy and gravel-cobble channel bars occur mostly in distances of 0–30 m from bank scours.

Erosion flood landforms and function of anthropogenic landforms during the flood

Erosion geomorphologic effects (or landforms) of the spring 2006 flood did not have such extend in floodplain of Sázava River as flood accumulation landforms. 176 bank scours in total length of 10 359 m were found. The average length of bank scours is 59 m and the median is 30 m. Bank scours reach the highest densities in the middle and upper part of the studied floodplain which means below Kácov, below Zruč nad Sázavou, below Ledec nad Sázavou, between Světlá nad Sázavou and Přibyslav (Fig. 7). If floodplain sediments were washed away somewhere, it always occurred on small areas related to narrowing of floodplain and mainly to surroundings of weirs (90% of all bank scours are located up to 10 m from weirs). These are places where bankfull stream reaches higher speed. This narrowing, accelerating formation of erosion furrows in floodplain, is of natural origin (e.g. alluvial fan nearby Chřenovice-Podhradí) or of anthropogenic character in areas of narrowing of cross-section profile of channel and floodplain (e.g. behind the bridge below Havlíčkův Brod) (Fig. 8).

Although there are settlements, roads, railways and technical buildings in Sázava River floodplain, it is possible to say that it kept its natural form in basic features. Low

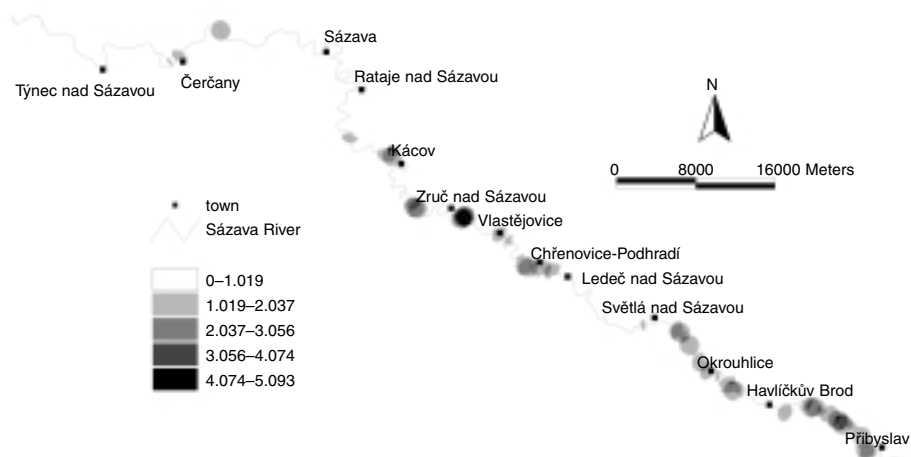


Fig. 7: Density of bank scours in Sázava River. Density represents number of bank scours per 1 km² for radius 1 000 m



Fig. 8: The bank scour and flood overbank accumulation behind a bridge below Havlíčkův Brod. The bridge and its road create a transverse barrier in Sázava River floodplain

level overbuilding of the Sázava River floodplain reflected in little destructive effects of the spring flood and relatively small material damages caused by the flooding. Roads and railways and anthropogenic mounds connected to them have mainly direction parallel to the flow and they are situated by border of the floodplain. Permanently inhabited settlements are in large extend connected to old alluvial terraces, it means outside inundation zone. Also huts and cottages are located by borders of the floodplain and on alluvial fans which rise up inundation areas. Bridges as potential barriers of flowing water have in most cases sufficient reserve for higher water-level stage of Sázava River.

Interpretation and discussion

Relations of fluvial accumulation landforms

Narrowing of cross-section profile, actually bottom of floodplain, reflects also in rise of kinetic energy of flowing water and in rise of erosive force of river. When floodplain widens again the speed decreases and so does transport ability of river which reflects in higher rate of accumulation. The character of Sázava River valley with its noticeable variation of width of valley bottom (especially in its middle and upper parts

of Sázava River) predisposed this river to “morphological” acceleration of stream speed. Exponential fall of flood overbank accumulations number depended on distance from channel of Sázava River (Fig. 3) corresponds to progress of decreasing sedimentation rate ($\text{g/cm}^2/\text{year}$) in five floodplains of rivers in England (Stour, Culm, Severn, Rother, Avon) (Welling, He, 1998). On the other hand, the flood on Sázava River does not belong to greater floods with respect to the progress of the relative amount of sedimentation in the flooded area (*sensu* Steiger, Gurnell, 2002). Proportionally higher representation of combined (coarser) flood overbank accumulations in bigger distances from the channel (25–75 m) can be explained by origin of these accumulations in areas with higher speed of bankfull flow which proves presence of coarse-grained material. Similar effect was described by Benedetti (2003) where the mean value of grain size of the 2001 flood deposits laying in the middle of McGregor Island was also greater than on its edge. While the flood went out, fine-grained sediments deposited in bigger distances than coarser sediments (e.g. Pizzuto, 1987; Asselmann, Middelkoop, 1995; Walling et al., 1997). It corresponds to spatial position of combined accumulations in areas of floodplain narrowing or by a mouth of stream tributary. Thus combined accumulations can be found in places with appropriate conditions for acceleration of bankfull stream. These parts of floodplain with combined accumulations can be considered as potential areas of more considerable effects of fluvial processes connected to overbank stream and following accelerated geomorphologic consequences such as creation of new erosion furrows, braiding of overbank stream and overbank sedimentation. It corresponds to findings of Pierce and King (2008), which show that the mean sediment overbank deposition in these areas (shoal and valley plug sites) is 2–10 times greater than at unchanneled sites.

Mutual communication between accumulation processes in the channel and floodplain of Sázava River proves dependency on spatial distribution of fluvial

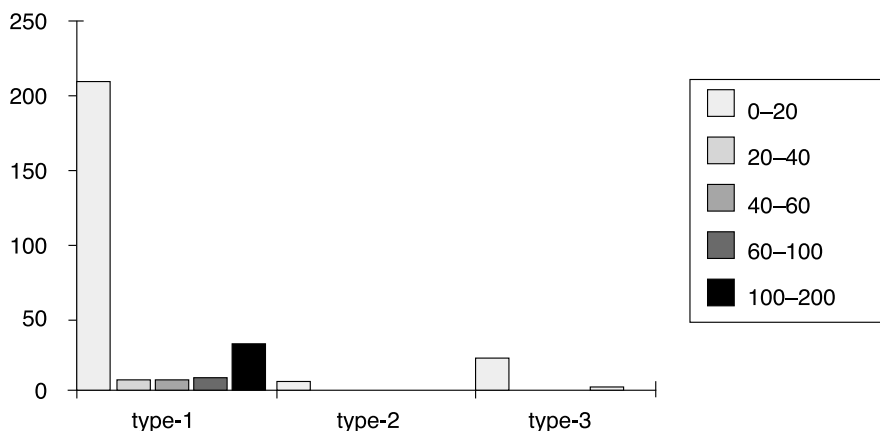


Fig. 9: Number of flood overbank accumulations according to type of sediment regarding distance from channel bars. Category values 0 – 30, 30 – 50 etc. represent distances (in meters) between channel bars and flood overbank accumulations of Sázava River. Type 1 – sandy flood accumulations, 2 – gravel-cobble flood accumulations, 3 – combined flood accumulations

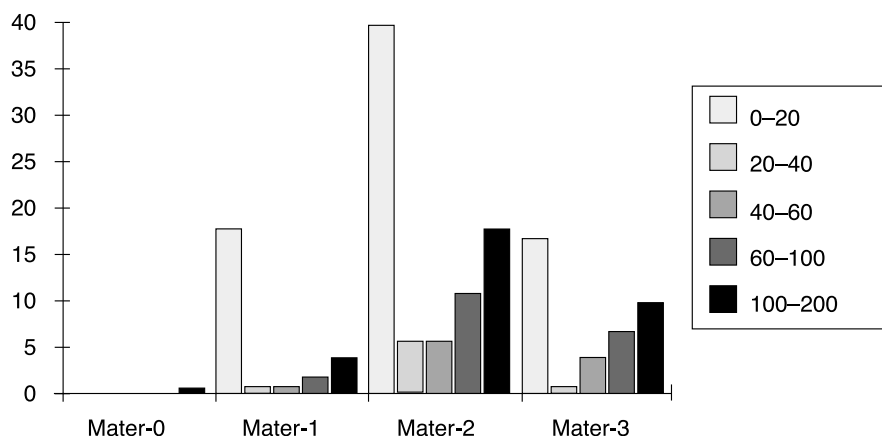


Fig. 10: Number of channel bars according to sediment type regarding distance from flood overbank accumulations. Category values 0–20, 20–40 etc. represent distances (in meters) between channel bars and flood accumulations in floodplain of Sázava River. Mater-0 ... non specified channel bars; Bars: Mater-1 ... sandy; Mater-2 ... gravel-cobble; Mater-3 ... combined

sediments in the area of the channel and its floodplain. 89% of flood overbank accumulations of the floodplain are located within 20 m distance from channel bars (Fig. 9). On the contrary up to distance of 20 m from flood accumulations in floodplain there are 50% of all channel bars (Fig. 10). This disproportion is caused by bigger spatial diffusion of flood overbank accumulations and by difference in number of channel bars (147) and flood overbank accumulations in the floodplain (303). Concerning distance characteristics we cannot forget that overvaluation of number of objects depending on increasing distance is caused by multiple count in of one object to greater distance categories. From the above mentioned it is clear that channel bars, which are more noticeable and can be mapped faster owing to their specific location, can be used for prediction of spatial distribution of flood accumulations in Sázava River floodplain. Thus, it is a simple method for predicting overbank deposition during floods with similar magnitude. Relationship between channel bars and flood overbank accumulations in the floodplain is determined by flood magnitude (Steiger, Gurnell, 2002) and it may not exist in great floods or it can be smaller because the maximum sedimentation might be expected to occur at the floodplain margins (Bridge, Leeder, 1979).

From the viewpoint of type of sediment it was proved that channel bars and flood accumulations in the floodplain have completely different grained composition (Fig. 11). This difference is statistically significant (it was tested by one-way ANOVA with current effect: $F(1, 448) = 189.35$, $p = 0.0000$). While channel bars have rather gravel-cobble character (56%), flood accumulations of floodplain are sandy (88%). It indicates different sedimentary conditions outside and inside the channel during the floods (also e.g. Steiger, Gurnell, 2002). It didn't significantly succeed to prove grain sizes equivalence between channel bar sediment and overbank sediment. Thus, type of flood overbank accumulation can not be estimated based on type of channel bars' sediment.

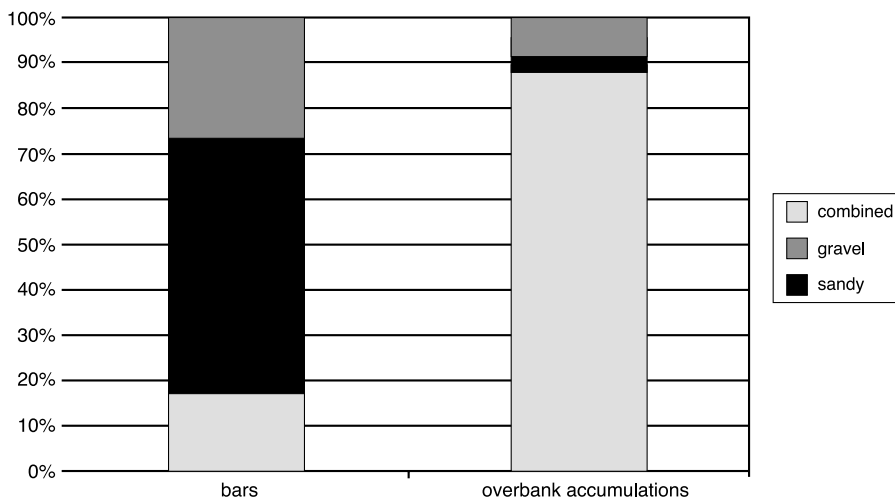


Fig. 11: Comparison of channel bars and flood overbank accumulations in floodplain according to sediment type

Analysis of bars' occurrence shows that they are concentrated in areas with lower sediment function of stream. Bends with abrupt change of stream direction, places under weirs and places near embouchure of adjacent river are these areas. Channel bars are created during floodings, actually at their end, and define places of higher rate of sediment deposition in the channel where they form potential barriers of the flow. It results in increased erosion activity and formation of bank scours. Thus, channel bars are also indicator of zones of higher erosion effectiveness of stream. On the other hand, this result reflects presence of weirs to which channel bars are connected. Weirs are places with greater sediment deposition and represent valley plugs in sense of Pierce and King (2008). Full-grown vegetation including self-seeding of shrubs and trees proves long-lasting process of sedimentation in chosen locality (channel bars). The results show that vegetation is more or less resistant regarding existing weirs and is developed in channel bars evenly along the whole stream without dependence on distance from the weirs.

From the viewpoint of spatial distribution of channel bars regarding weirs, mainly of gravel-cobble and combined channel bars, we can say that weirs influence change of behavior of Sázava River (from the viewpoint of sedimentation) in distance of approximately 60 m. Rapid decrease of stream energy proves occurrence of channel bars with normal gradation whose presence out of areas influenced by weirs is very low. On the contrary, channel bars without gradation are typical for parts of Sázava River which are not influenced directly by weirs, i.e. in distances bigger than 100 m. More frequent occurrence of channel bars placed under weirs in the centre of the channel is caused by the fact that the channel is wider in these areas, that way larger islands of channel bars can be created. It follows that smaller channel bars connected to banks are typical for parts of Sázava River which are not influenced by direct impact of weir's steps. Zielinski (2003) showed that the central longitudinal bars were formed in straight flow of Nysa River. Thus, this type of channel bars is typical for mid-channel streamline.

Bank scours bond flood overbank accumulations in floodplain (Fig. 12) and channel bars (Fig. 13) to themselves. 80% of all flood accumulations in floodplain are located up to 10 m from bank scours. All gravel-cobble and combined flood overbank accumulations in floodplain are located up to 10 m from bank scours. 83% of all bars are located up to 10 m from bank scours. Sedimentation occurs when ability of stream drifting lowers. It is related to speed change of water flowing in the channel (with lower friction) and in the floodplain (with higher friction) during flood (Bridge, 2003). This change of speed of stream is the biggest between the channel and the rest of the floodplain. That is why the biggest sedimentation and the most thick flood accumulations are located tightly by banks of floodplain (e.g. Walling, He, 1998; Benedetti, 2003). It shows also position of well transportable sandy flood accumulations (Fig. 12) regarding bank scours which line banks of the channel. Many flood overbank accumulations are the largest in immediate vicinity of channel. Then number of flood overbank accumulations decreases slowly in distance of 10–20 m from the midstream of channel and in distance of 20–50 m from the channel midstream it has an increasing tendency which is caused probably by sedimentation of material from damaged nearby bank. It correlates with a conceptual model of riparian sedimentation patterns – type (b): sedimentation pattern generated by a larger flood within a geomorphologically simple and constrained setting (Steiger, Gurell, 2002).

Weirs in period of time of high water level are potential places of stream instability and cause significant changes of behavior of this stream. Noticeable relation effects of fluvial lateral erosion and origin of bank scours in their vicinity proves that.

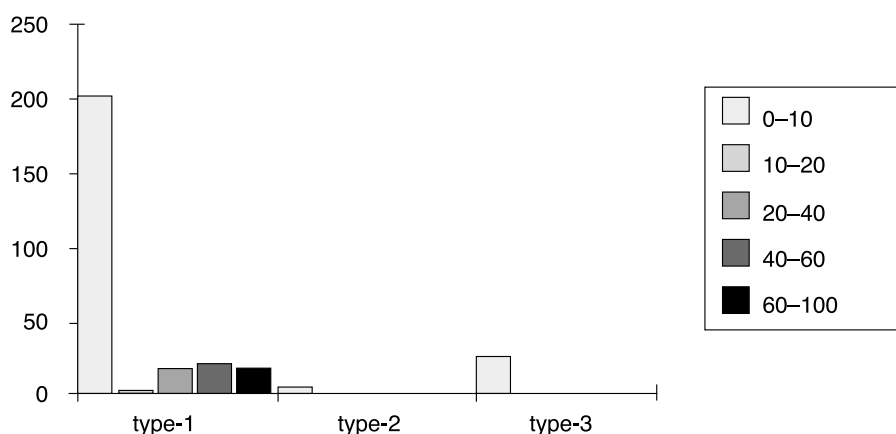


Fig. 12: Number of flood overbank accumulations in floodplain by type of sediment regarding distance from bank scours. Category values 0–10, 10–20 etc. represent distances (in meters) between overbank accumulations of Sázava River and bank scours. Type 1 – sandy flood accumulations, 2 – gravel-cobble flood accumulations, 3 – combined flood accumulations

Conclusion

Despite the spring flood in Sázava River caused inundation of the whole floodplain and in some places crossed its border, its consequences were not as serious as in streams in river basin Otava after summer flooding in 2002 (Křížek, Engel, 2006). It is mostly caused by small build over of floodplain of Sázava River and its channel. In spite of this some anthropogenic impacts in the floodplain and channel of Sázava River represent potential risks of overbank of stream which bond accelerated erosion and accumulation. It concerns mainly weirs.

During the spring flood in 2006 fluvial accumulation processes prevailed over erosion. It turned out that places with the highest number of flood overbank accumulations in floodplain agree with places with the greatest occurrence of flood overbank accumulations in the floodplain. Number of flood overbank accumulations falls exponentially with increasing distance from the river channel (Fig. 13). Proportionally higher representation of combined (coarser) flood overbank accumulations in bigger distances from the channel (25–75 m) can be explained by origin of these accumulations in areas of higher speed of bankfull flow which proves presence of coarse-grained material. Benedetti (2003) also described a similar rise of grain size depending on distance from the channel. Spatial distribution of channel bars, flood overbank accumulations in floodplain and bank scours is uneven and significantly connected to weirs.

Channel bars are a good typical feature of sedimentary activity of stream. Channel bars are connected to weirs whose immediate effective sedimentary activity reaches in Sázava River to distance up to 60 m. Smaller channel bars connected to banks are typical for parts of river which are not influenced by weirs. Channel bars also have close connection with deposition of flood accumulations within floodplain (89% of flood accumulations in floodplain are located up to 20 m from a channel bar) and with spatial distribution of bank scours. Channel bars

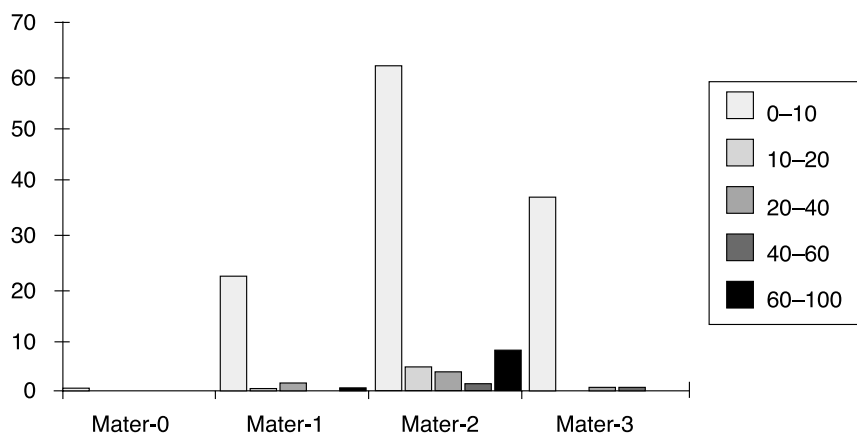


Fig. 13: Number of channel bars by type of sediment regarding distance from bank scours. Category values 0 – 10, 10 – 20 etc. represent distances (in meters) between channel bars and bank scours. Mater-0 ... non specified channel bars; Bars: Mater-1 ... sandy; Mater-2 ... gravel-cobble; Mater-3 ... combined

can be used for prediction of higher accumulation activity occurrence in floodplain but it can not be used for prediction of sedimentation type. Combined flood overbank accumulations in floodplain are concentrated to places with multiple flood sedimentation and they support areas of floodplain which are more endangered by geomorphologic features of flood.

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Résumé

Prostorové rozmístění erozních a akumulačních fluválních tvarů na Sázavě po povodni 2006

Předložený článek pojednává o korytových akumulacích, povodňových akumulacích údolní nivy a břehových nátržích jako nejdůležitějších geomorfologických projevech jarní povodně na Sázavě v roce 2006. Zvláště pak si všímá jejich prostorového rozmístění vzhledem ke korytu, jezům a sobě navzájem. Je zkoumán vztah mezi základními sedimentologickými charakteristikami těchto korytových akumulací, jejich rozmístěním a polohou vzhledem k jezům.

Problematika akumulace v údolní nivě je řešena v zahraničí pomocí analýzy přírůstu sedimentů, např. Nicholas and Walling (1997), Walling and He (1998), Steiger and Gurnell (2001), přičemž se autoři snaží pomocí svých výsledků predikovat ohrožená místa v příčném profilu údolních niv. Výhodou takového přístupu je, že lze pozorovat v jednom místě několik sledů povodní. Nevýhodou je však, že u starých povodní se neví, jak patřičná část nivy vypadala, např. kudy vedlo koryto, což je zásadní pro distribuci sedimentů. V Čechách a na Slovensku převažují na toto téma většinou práce popisné bez statisticky podložených zobecnění.

Během jarního období 2006 došlo k tání sněhu na velkých plochách středních poloh, což vedlo k rozvodnění velkého množství řek a následným povodním. 20. 3. 2006 činily celkové zásoby vody ve sněhové pokrývce v celém povodí Sázavy 446,3 mil. m³ (Šercl et al., 2006). Dvouvrcholová povodeň kulminovala 30. 3. 2006 a maximální průtok Sázavy v Nespekách byl 536 m³/s, pro srovnání průměrný roční průtok Sázavy v Nespekách je 23,4 m³/s (Šercl et al., 2006).

Přestože jarní povodeň na Sázavě způsobila zaplavení celé údolní nivy a místy překročila i její hranici, nebyly její následky takové jako na vodních tocích povodí Otavy po letních povodních v roce 2002 (Křížek, Engel 2006). Z velké míry je to dáno menším zastavěním údolní nivy Sázavy a jejího koryta. Přesto však některé antropogenní zásahy do údolní nivy a koryta Sázavy představují potenciální rizika v případech vyběžení vodního toku, které na sebe váží akceleraovanou erozi a akumulaci. Jedná se zejména o jezy.

Během jarní povodně 2006 převládaly v údolní nivě Sázavy akumulační fluvialní procesy nad erozními. Ukázalo se, že místa s největším počtem povodňových akumulací v údolní nivě mimo koryto se shodují s místy výskytu nejrozhlednějších mimokorytových povodňových akumulací. Rozmístění korytových akumulací, povodňových akumulací údolní nivy a břehových nátrží je nerovnoměrné a značně vázané na jezy.

Exponenciální pokles množství mimokorytových povodňových akumulací se vzdáleností od koryta Sázavy (obr. 3) odpovídá průběhu poklesu sedimentační rychlosti ($\text{g}/\text{cm}^2/\text{rok}$), podobně jak to bylo zjištěno v údolních nivách anglických řek (Welling, He, 1998). Poměrově větší zastoupení kombinovaných (hrubo-zrnějších) mimokorytových povodňových akumulací ve větších vzdálenostech od koryta 25–75 m je možné vysvětlit vyšší rychlostí vyběženého vodního toku v této zóně, což dokládá i přítomnost hrubozrnějšího materiálu v této zóně. Podobný nárůst zrnitosti od koryta popisuje i Benedetti (2003).

Korytové akumulace jsou dobrým charakteristickým znakem sedimentační činnosti vodního toku. Korytové akumulace jsou vázány na jezy, jejichž bezprostřední efektivní sedimentační působnost zasahuje na Sázavě do vzdálenosti 60 metrů. Menší korytové akumulace spojené se břehy jsou typické pro části toku, které nejsou ovlivněny jezy. Stejně tak mají korytové akumulace těsnou souvislost s ukládáním povodňových akumulací v rámci údolní nivy (89% povodňových akumulací údolní nivy se nachází do vzdálenosti 20 m od korytové akumulace) a s rozmístěním břehových nátrží. Korytové akumulace lze v povodni podobného rozsahu použít při predikci výskytu zvýšené akumulační činnosti v údolní nivě, avšak nelze ji využít z hlediska predikce typu sedimentace. Zde je třeba připomenout, že síla vazby mezi korytovými akumulacemi a povodňovými akumulacemi v údolní nivě je ovlivněna magnitudem povodně (sensu Steiger, Gurnell, 2002), s tím, že u velkých povodní nemusí vzájemná podmíněnost existovat a nebo může být menší, neboť maximum sedimentace může být v takových případech lokalizováno do okrajů údolní nivy (Bridge, Leeder, 1979).

Kombinované povodňové akumulace v nivě jsou soustředěny do míst s vícenásobnou povodňovou sedimentací a dokumentují ta místa nivy, která jsou více ohrožena geomorfologickými projevy povodní.

Obr. 1: Studované území.

Obr. 2: Hustota fluvialních povodňových akumulací v údolní nivě. Hustota udává počet fluvialních akumulací na 1 km^2 pro rádius 1 000 m. A – každá mimokorytová povodňová akumulace je reprezentována polohou těžiště. B – každá mimokorytová povodňová akumulace je reprezentována sítí bodů, jejichž počet je vážen velikostí příslušného akumulačního tvaru.

Obr. 3: Poměrné zastoupení jednotlivých typů fluvialních akumulací vzhledem ke vzdálenosti od koryta (střednice vodního toku). Type 1 – jsou hlinitopísčité fluvialní akumulace, 2 – šterkovokamenité fluvialní akumulace, 3 – kombinované fluvialní akumulace. Kategorizační hodnoty, 0–25, 25–50, atd. představují vzdálenosti v metrech od střednice vodního toku.

Obr. 4: Hustota korytových akumulací na Sázavě. Hustota udává počet fluvialních akumulací na 1 km^2 pro rádius 1 000 m.

Obr. 5: Rozmístění korytových akumulací vzhledem ke vzdálenosti od jezů (v metrech). Mater-0 ... nespecifikované korytové akumulace; Korytové akumulace: Mater-1 ... hlinitopísčité; Mater-2 ... šterkovokamenité; Mater-3 ... kombinované.

Obr. 6: Korytová akumulace pokrytá travou pod jezem u Chřenovic.

Obr. 7: Hustota břehových nátrží na Sázavě. Hustota udává počet břehových nátrží na 1 km^2 pro rádius 1 000 m.

Obr. 8: Břehová nátrž a fluvialní akumulace v údolní nivě za mostem pod Havlíčkovým Brodem. Most a nájezd na něj působí jako příčná překážka v údolní nivě Sázavy. Na fotografii je patrná trasa proudění vyběžené Sázavy.

Obr. 9: Zastoupení mimokorytových povodňových akumulací údolní nivy dle typu sedimentu vzhledem ke vzdálenosti od korytových akumulací. Kategorizační hodnoty, 0–30, 30–50, atd. představují vzdálenosti v metrech mezi akumulacemi v korytě a mimo koryto. Type 1 – jsou hlinitopísčité fluvialní akumulace, 2 – šterkovokamenité fluvialní akumulace, 3 – kombinované fluvialní akumulace.

Obr. 10: Zastoupení korytových akumulací dle typu sedimentu vzhledem ke vzdálenosti od mimokorytových povodňových akumulací údolní nivy. Kategorizační hodnoty, 0–30, 30–50, atd. představují vzdálenosti

v metrech mezi akumulacemi v korytě a mimo koryto. Mater-0 ... nespecifikované korytové akumulace; Korytové akumulace: Mater-1 ... hlinitopísčité; Mater-2 ... štěrkovokamenité; Mater-3 ... kombinované.

Obr. 11: Srovnání korytových akumulací a povodňových mimokorytových akumulací údolní nivy z hlediska typu sedimentů.

Obr. 12: Počet mimokorytových povodňových akumulací údolní nivy dle typu sedimentu vzhledem ke vzdálenosti od břehových nátrží. Kategorizační hodnoty, 0–10, 10–20, atd. představují vzdálenosti v metrech mezi mimokorytovými akumulacemi v údolní nivě Sázavy a břehovými nátržemi. Type 1 – jsou hlinitopísčité fluvialní akumulace, 2 – štěrkovokamenité fluvialní akumulace, 3 – kombinované fluvialní akumulace.

Obr. 13: Počet korytových akumulací dle typu sedimentu vzhledem ke vzdálenosti od břehových nátrží. Kategorizační hodnoty, 0–10, 10–20, atd. představují vzdálenosti v metrech mezi korytovými akumulacemi a břehovými nátržemi. Mater-0 ... nespecifikované korytové akumulace; Korytové akumulace: Mater-1 ... hlinitopísčité; Mater-2 ... štěrkovokamenité; Mater-3 ... kombinované.

Tab. 1: Zjišťované charakteristiky vybraných tvarů při mapování geomorfologických následků jarní povodně 2006 na Sázavě.

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